

## IUPAC Task Group on Atmospheric Chemical Kinetic Data Evaluation – Data Sheet VI.A4.26 HET\_SL\_26

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### HOBr + HCl (aqueous sulphuric acid aerosol) → BrCl + H<sub>2</sub>O

#### Experimental data

Parameter	wt% H <sub>2</sub> SO <sub>4</sub>	Temp./K	Reference	Technique/ Comments
$\gamma(\text{HOBr})$				
0.023	69.8% (HCl = $0.22 \times 10^{12} \text{ cm}^{-3}$ )	228	Abbatt, 1995	CWFT-MS (a)
0.060	(HCl = $0.96 \times 10^{12} \text{ cm}^{-3}$ )			
0.12	(HCl = $2.5 \times 10^{12} \text{ cm}^{-3}$ )			
0.20	(HCl = $6.0 \times 10^{12} \text{ cm}^{-3}$ )			
0.2	60% ( p(HCl) = $2.7 \times 10^{-7}$ mbar)	210	Hanson & Ravishankara, 1995	CWFT-MS (b)
$\gamma(\text{HCl})$				
0.11±0.04	59.7% ( p(HOBr) = $8 \times 10^{-7}$ mbar)	228	Waschewsky and Abbatt, 1999	CWFT-MS (c)
0.032±0.008	65.6% ( p(HOBr) = $1 \times 10^{-6}$ mbar)	228		
0.060±0.015	70.1% ( p(HOBr) = $3.4 \times 10^{-5}$ mbar)	228		
0.19	58%; ( p(HOBr) = $1 \times 10^{-5}$ mbar)	250	Hanson, 2003	CWFT-CIMS (d)
0.07				
0.011	62%; ( p(HOBr) = $1 \times 10^{-5}$ mbar)			
0.002	65%; ( p(HOBr) = $1 \times 10^{-5}$ mbar)			
	69.5%; ( p(HOBr) = $1 \times 10^{-5}$ mbar)			

#### Comments

- (a) HOBr ( $<1 \times 10^{12}$  molecule  $\text{cm}^{-3}$ ) was generated by the reaction sequence:  $\text{H} + \text{NO}_2 \rightarrow \text{OH} + \text{NO}$  and  $\text{OH} + \text{Br}_2 \rightarrow \text{HOBr}$ , ionised by electron impact and detected as  $\text{HOBr}^+$ . The carrier gas flow was humidified to maintain the H<sub>2</sub>SO<sub>4</sub> concentration (69.8 wt %). Excess HCl (gas phase concentration  $(0.2\text{-}7.0) \times 10^{12}$  molecule  $\text{cm}^{-3}$ ) added simultaneously with HOBr ( $>1 \times 10^{12}$  molecule  $\text{cm}^{-3}$ ). HOBr loss was first order, and no systematic dependence of  $\gamma$  on [HOBr] was observed. Using measured values of  $HD_1^{1/2}$  for HOBr ( $(20 \pm 10) \text{ M atm}^{-1} \text{ cm s}^{-1/2}$ ) and for H\* for HCl the cited  $\gamma$  value constrains the bimolecular rate constant for the HOBr + HCl reaction to be  $1.4 \times 10^5 \text{ M}^{-1} \text{ s}^{-1}$ .
- (b) HOBr was generated from the hydrolysis reaction of BrONO<sub>2</sub> on 60% H<sub>2</sub>SO<sub>4</sub> and detected using SF<sub>6</sub><sup>-</sup> chemi-ions. BrCl was observed as a gas phase product indicating a fast effective bimolecular reaction on H<sub>2</sub>SO<sub>4</sub>. A bimolecular rate constant for the HOBr + HCl reaction of  $(0.3\text{-}1) \times 10^5 \text{ M}^{-1} \text{ s}^{-1}$  was derived from the uptake data.

- (c) HOBr ( $\approx 5 \times 10^{-10}$  bar, measured by UV absorption at 254 nm) was generated ex-situ by passing a humidified flow of Br<sub>2</sub> in He over HgO, ionised by electron impact and detected as HOBr<sup>+</sup>. Decay of HCl and formation of BrCl was measured in presence of excess HOBr. Uptake coefficients determined from both kinetic curves were within 10%.  $\gamma$  increased with  $p(\text{HOBr})^{1/2}$  at 65-70%. Bimolecular rate constants for the HOBr + HCl reaction derived from the uptake data using  $H^*(\text{HOBr})$  determined in the same study were in the range of  $(1.6 - 270) \times 10^5 \text{ M}^{-1} \text{ s}^{-1}$ , increasing strongly with wt% H<sub>2</sub>SO<sub>4</sub> and less so with temperature (213 -228 K). The [H<sub>2</sub>SO<sub>4</sub>] dependence is attributed to an acid catalysed mechanism and the temperature dependence to diffusion limitation. The following expression is given for calculations of  $\gamma$  for stratospheric conditions:  $k^{\text{II}} = \exp(0.542 \times (\text{wt}\% \text{ H}_2\text{SO}_4) - 6.44 \times 10^3/T + 10.3)$
- (d) Rotating CWFT with stirring of the H<sub>2</sub>SO<sub>4</sub> film (58 – 70 wt. %). HOBr (typically  $10^{-10}$  atm) was generated by reacting BrONO<sub>2</sub> with water and detected using SF<sub>6</sub><sup>-</sup> chemi-ions. First order loss of HCl was measured in the presence of excess HOBr  $(0.2\text{-}2.0) \times 10^{-10}$  bar.  $\Gamma_b$  was extracted assuming  $\alpha = 1.0$ , and was  $\propto p(\text{HOBr})^{1/2}$ . Using  $H_{\text{HCl}}$  from Carslaw et.al. (1995) and  $H_{\text{HOBr}}$  determined in the same study (see reaction VI.A4.16) values of  $k^{\text{II}}$  were derived. The values ( $k^{\text{II}} = (2.5\text{-}6.0) \times 10^7 \text{ M}^{-1} \text{ s}^{-1}$  over the range 58 -69.5% H<sub>2</sub>SO<sub>4</sub> at 250 K) showed a lower dependence on acid strength compared to results of Waschewsky, G. C. G. and Abbatt, J. P. D., (1999), obtained at lower temperatures.

### Preferred Values

Parameter	Value	T/K
$\alpha_s$	1.0	190 - 250
$H^*_{\text{HOBr}}/\text{M.bar}^{-1}$	$5.22 \times 10^{-5} \exp(5427/T)$	210 - 270
$H^*_{\text{HCl}}/\text{M.bar}^{-1}$	$(0.094 - 0.61X + 1.2X^2) \exp(-8.68 + (8515 - 10718X^{0.7})/T)$	203
$k^{\text{II}}(\text{M}^{-1}\text{s}^{-1})$	$\exp(154 - 1.63.W) \exp(-(38500 - 478.W)/T)$	200-230
$D_{\text{HCl}}/\text{cm}^2\text{s}^{-1}$	$7.8 \times 10^{-8} T/\eta$	190-240
<i>W</i> = wt% H <sub>2</sub> SO <sub>4</sub> , <i>X</i> = mole fraction H <sub>2</sub> SO <sub>4</sub>		
<i>Reliability</i>		
$\Delta \log(\alpha)$	0.3	298

### Comments on Preferred Values

The experimental studies show that the uptake coefficient of HOBr or HCl on H<sub>2</sub>SO<sub>4</sub>/H<sub>2</sub>O containing HCl or HOBr increases with the concentration of the co-reactant in solution. The  $\gamma_{\text{HOBr}}$  values from the earlier study of Abbatt (1995) are uncertain because the assumed relative solubilities of HOCl and HCl were incorrect. The measured  $\gamma_{\text{HCl}}$  in the studies of Waschewsky and Abbatt (1999) and Hanson (2003) agree well and the  $\gamma$  values show a strong increase with the water content of the H<sub>2</sub>SO<sub>4</sub> and a rather weak temperature dependence. The kinetics are consistent with the resistance model with a large surface accommodation coefficient ( $\alpha_s = 1$ ), and rate of uptake controlled by bulk phase chemical reaction HOBr+HCl:

$$\frac{1}{\gamma} = \frac{1}{\alpha_s} + \frac{1}{\Gamma_b} \quad (1)$$

$$\Gamma_b = \frac{4RTH_{HOBr}}{c} \sqrt{D_l k^{II} H_{HCl}^* p_{HCl}} \quad \text{for } \gamma_{HOBr} \quad (2)$$

$$\text{or } \Gamma_b = \frac{4RTH_{HCl}^*}{c} \sqrt{D_l k^{II} H_{HOBr}^* p_{HOBr}} \quad \text{for } \gamma_{HCl} \quad (3)$$

where  $H^*$  are the Henry's solubility constants and  $D_l$  are the liquid diffusion coefficients for HOBr and HCl which are a function of mole fraction of  $H_2SO_4$ ,  $X$ .

There are significant differences in the heterogeneous reaction rate constants derived from the experimental uptake coefficients in the two studies. Thus the values of  $k^{II}$  derived by Waschewsky and Abbatt, (1999) from their data at 213 - 238 K was a factor of  $\sim 8$  lower than the values expected from extrapolation from the 250 K value of  $k^{II}$  derived by Hanson (2003), based on an estimated activation energy of 15 kJ mol<sup>-1</sup>. This discrepancy can be largely attributed to the use of different values for  $H(HOBr)$  in the derivation of  $k^{II}$ . The preferred value for  $H(HOBr) = 5.22 \times 10^{-5} \exp(5427/T)$  is based on Hanson (2003) (see IUPAC evaluation VI.4.16), which showed only a weak dependence on acid strength in the range 60 - 70%  $H_2SO_4$ . Using this choice of  $H(HOBr)$  Hanson gave coherent set of  $k^{II}$  as a function of T and wt%  $H_2SO_4$ , from the results of all studies.  $k^{II}$  showed a very strong T-dependence at 60 wt% acid and a much weaker one at 70%. Hanson gives an expression for  $k^{II}$  for stratospheric conditions ( $W = \text{wt}\% H_2SO_4$ ):

$$k^{II} (\text{M}^{-1}\text{s}^{-1}) = \exp(154 - 1.63.W) \exp(-(38500 - 478.W)/T)$$

This equation reproduces the experimentally derived  $k^{II}$  values at 210 – 228 K but the predicted values at 238 and 250 K are over estimated by a factor of 3 and 6 respectively. Nevertheless this is the recommended expression to calculate the values of the uptake coefficients for stratospheric conditions, using eqs.(1)–(3). The  $X$ -dependent expressions used for  $H_{HCl}^*$  (based on Shi et al, 2001; see this evaluation: VI.A4.14) and for  $H_{HOBr}^*$  (based on Hanson, 2003; see this evaluation: VI.A4.16), are given in the Table of preferred values. The expression for  $D_l(HCl)$  from Klassen et al., 1998, based on viscosity data given by Shi et al., 2001 (see this evaluation: VI.A4.25).

Figure 1 shows a comparison of calculated  $\gamma_{HCl}$  with selected experimental values grouped for acid strengths of  $\sim 59$ , 65 and 70%. The prediction is only good for lower acid strengths and temperatures.

## References

- Abbatt, J. P. D.: J. Geophys. Res. 100, 14009-14017, 1995.  
Hanson, D.R. and Ravishankara, A.R.: Geophys. Res. Lett. 22, 385 (1995).  
Hanson, D.R.: J. Phys. Chem. A102, 4794 (1998)  
Hanson, D. R.: J. Geophys. Res. 108, doi:10.1029/2002JD002519, 2003.  
Klassen, J. K., Hu, Z. and Williams, L. R.: J. Geophys. Res. 103, 16197, 1998.  
Shi, Q., Jayne, J. T., Kolb, C. E., Worsnop, D. R., and Davidovits, P.: J. Geophys. Res., 106, 24259, 2001.  
Waschewsky, G. C. G. and Abbatt, J. P. D.: J. Phys. Chem. A 103, 5312 1999.

**Figure 1:** Plot of  $p(\text{HOBr})$  dependence of uptake coefficient  $\gamma(\text{HCl})$  at different  $\text{H}_2\text{SO}_4$  wt%. symbols show experimental data: open symbols = 70%; half filled symbols = 65%; closed symbols = 59%. Lines show predicted uptake coefficients using recommended parameters. Temperature = 213 - 250 K.

