

IUPAC Task Group on Atmospheric Chemical Kinetic Data Evaluation

Data Sheet VI.A1.20 HET_H2OL_20

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This data sheet evaluated: June 2017; last change in preferred values: June 2017.

HO₂ + H₂O (aq) → products

Experimental data

	Temp./K	Reference	Technique/ Comments
<i>Accommodation coefficients: α</i>			
> 0.01	275	Hanson et al, 1992	WWFT-LIF (a)

Comments

- (a) Uptake of HO₂ ($5\text{-}30 \times 10^{10}$ molecule cm⁻³) to a film (0.2 mm thick) of de-ionised water, or water containing 10⁻³ M CuSO₄. HO₂ was formed in the reaction of F with H₂O₂ and detected as OH after reaction with NO. HO₂ uptake was limited by diffusion through the 1 Torr of He bath gas. Levels of HO₂ were sufficiently low to neglect loss due to gas-phase self-reaction. Addition of CuSO₄ had no effect on HO₂ loss rates.

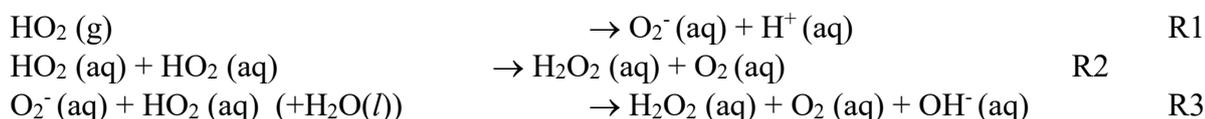
Preferred Values

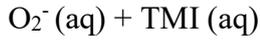
Parameter	Value	T/K
α_b	> 0.2	270-300
Reliability $\Delta(\log \alpha_b)$	undetermined	

Comments on Preferred Values

The sole experimental study of the uptake of HO₂ to pure water returned a lower limit to the accommodation coefficient of $\alpha_b > 0.01$. This is consistent with molecular dynamics calculations of the HO₂-water interaction (Morita et al., 2004), which suggest that the accommodation coefficient could be unity, and also with more sensitive experiments on other aqueous surfaces such as (NH₄)₂SO₄ and NH₄HSO₄ (see datasheets VI.A3.09 and VI.A3.10). We thus prefer a value of α_b of > 0.2 in order to have an internally consistent recommendation for all aqueous substrates.

The uptake of HO₂ in aqueous solution is presently believed to be driven by self-reaction and acid-base dissociation of HO₂ (pK_a ~ 4.7) with formation of H₂O₂ (R2, R3). In the presence of transition metal ions (TMI) the reaction of HO₂ and especially O₂⁻ (R4) can be important:





R4

If a first-order loss process for HO₂ or O₂⁻ in the aqueous phase dominates (e.g. reaction with TMI such as Cu(II)), the uptake coefficient can be calculated from the expression below:

$$\frac{1}{\gamma} = \frac{1}{\alpha_b} + \frac{1}{\Gamma_{\text{TMI}}}$$

$$\frac{1}{\Gamma_{\text{TMI}}} = \frac{c}{4H^{\text{eff}} RT \sqrt{D_1 k_{\text{TMI}} [\text{coth}(r_p / l_{\text{rd}}) - (l_{\text{rd}} / r_p)]}} \quad l_{\text{rd}} = \sqrt{\frac{D_1}{k_{\text{TMI}} [\text{TMI}]}} \quad (1)$$

$H^{\text{eff}} = H^{\text{HO}_2} (1 + K_{\text{eq}} / [\text{H}^+])$, $K_{\text{eq}} = 2.1 \times 10^{-5} \text{ M}$ at 298 K (Jacob, 2000) and $H^{\text{HO}_2} = 9.5 \times 10^{-6} \exp(5910/T) \text{ M atm}^{-1}$ (Hanson et al., 1992). k_{TMI} is the second order rate coefficient for the reaction of HO₂ and O₂⁻ with transition metal ions. $D_1 = \{1 \times 10^{-5} (T/298)\} / (1.09 \times 10^8 \exp(-0.068T) + 0.873) \text{ cm}^2 \text{ s}^{-1}$ (Schwartz, 1984; Thornton et al., 2008) where the denominator in the D_1 term was derived from a fit to the water viscosity data of Hallett (1963). The size dependent correction factor, with r_p denoting the particle radius and l_{rd} the reacto-diffusive length, assures proper representation when the kinetic regime changes from reaction-diffusion towards volume limited kinetics at low TMI concentrations. As discussed by Thornton et al. (2008) and Mozurkewich et al. (1987), the apparent first order loss rate coefficient of HO₂ / O₂⁻ in the presence of TMI is around three orders of magnitude lower than based on the second order rate coefficient for dilute aqueous solutions (Bielski et al., 1985) possibly due to high ionic strength effects. Therefore, the value recommended for deliquesced organic and inorganic aqueous aerosols is most likely too low for dilute aqueous solutions. We therefore do not provide a recommendation for k_{TMI} for this case. Further heterogeneous kinetics experiments with variation of pH, [Cu(II)] and use of other TMI are needed to constrain γ for dilute aqueous solutions.

According to the reaction scheme (R1-R4) above, in the absence of TMI, the rates of loss of aqueous-phase HO₂ are quadratically dependent on [HO₂]_{aq} and [O₂⁻]_{aq} and are thus strongly dependent on the gas-phase concentration of HO₂. At low, atmospherically relevant, HO₂ concentrations the liquid phase reactions become rate limiting and γ is expected to be much smaller, as observed in dilute solutions by Mozurkewich et al. (1987) and also in the George et al. (2013) and Lakey et al. (2016) data for (NH₄)₂SO₄ and for aqueous organic aerosols. Thornton and Abbatt (2005) suggest that the rate of loss of HO₂ from the gas-phase (in molecule cm⁻³ s⁻¹) is best described by a system in thermodynamic (Henry's law) equilibrium so that (Thornton et al., 2008):

$$\frac{1}{\gamma} = \frac{1}{\alpha_b} + \frac{1}{\Gamma_{\text{self}}} \quad \frac{1}{\Gamma_{\text{self}}} = \frac{3cN_{\text{Av}}}{8000(H^{\text{eff}} RT)k_{\text{aq}}[\text{HO}_2]r_p} \quad (2)$$

k_{aq} can be calculated from the rate coefficients for R2 (k_2) and R3 (k_3) (Bielski et al., 1985) and the pH:

$$k_{\text{aq}} = \frac{k_2 + \left(\frac{K_{\text{eq}}}{[\text{H}^+]_{\text{aq}}} \right) k_3}{\left(1 + \frac{K_{\text{eq}}}{[\text{H}^+]_{\text{aq}}} \right)^2} \quad (3)$$

Equations (1) and (2) consistently describe the uptake of HO₂ in the presence and absence of transmission metal ions. At low HO₂ concentrations, the overall γ can be obtained by adding the corresponding resistors:

$$\frac{1}{\gamma} = \frac{1}{\alpha_b} + \frac{1}{\Gamma_{\text{TMI}} + \Gamma_{\text{self}}} \quad (4)$$

Equation (4) explains the measured dependence of γ over the full range of Cu(II) concentrations, with the caveats for the rate coefficients for dilute aqueous solutions mentioned above. As discussed by Hanson et al. (1992) and Thornton and Abbatt (2005), the parameterization suggested here is very sensitive to the solubility of HO₂ (H^{HO_2}), its temperature dependence and on the aerosol pH. We refer to recent publications for a more detailed description of the effect of different parameterisation schemes (Thornton et al., 2008; Macintyre and Evans, 2011).

References

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